

Genesis and spatial extension of a 1000m deep convection event in the Irminger Sea in 2011-2012

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The logo for Ifremer, featuring a grey silhouette of a fish jumping out of a yellow horizontal bar. The word "Ifremer" is written in black text on the yellow bar.

Ifremer



METEO FRANCE
Toujours un temps d'avance

Summary

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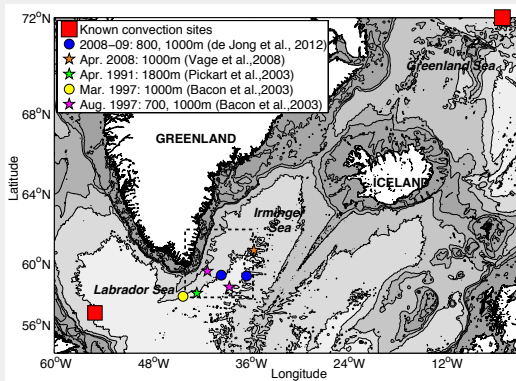
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Bibliographic context



Convection sites in the North Atlantic Ocean and observations of deep mixed layers in the Irminger Sea

Past observations of convection in the Irminger Sea are **limited in space and time**

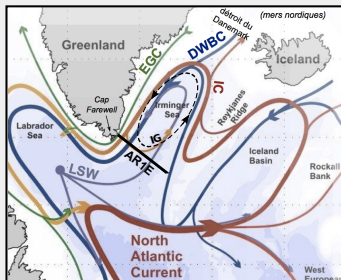
- ✧ Lack of data, especially during wintertime
- ✧ Focus on the Labrador site



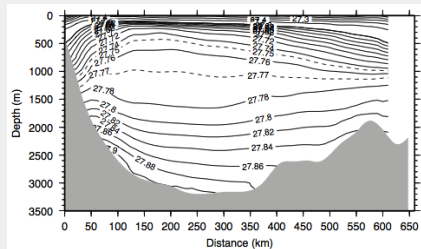
Mechanisms of convection in the Irminger Sea

1. Preconditioning

- Cyclonic circulation : Irminger Gyre [Lavender et al., 2000]
- Doming of isopycnals [Pickart et al., 2003]
- LSW in the Irminger intermediate layers [Pickart et al., 2003]



Irminger Sea circulation [P. Lherminier, LPO]



Vertical section (AR1E) of potential density

[Pickart et al., 2003]



Mechanisms of convection in the Irminger Sea

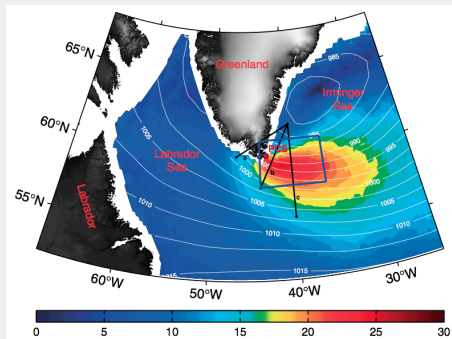
2. Atmospheric forcings

Greenland relief influence : formation of Greenland Tip Jet events

Greenland Tip Jets :

- regional-scale atmospheric events
- high wind speed (westerly)
- duration < 1 day
- induce heat loss that can reach 1000 W.m^{-2} locally

Mean (1999-2002) Tip Jet QuikSCAT winds (m.s^{-1}) and ERA 40 sea level pressure (hPa) [Vage et al., 2009]; Blue box : 'TJ box'



Questions

ARGO data, sampling the Irminger Sea all year (especially during winters) give informations about the ocean AT BASIN SCALE, DURING the convection event.



ARGO data permit to identify a deep convection event during the 2011-2012 winter.

- What is the spatial extent of deep convection in the Irminger Sea ?
- Can we identify the sequence of atmospheric forcings responsible for convection ?
- Can we use the 2011-2012 event to have a better understanding of the past events that could not be observed because of the lack of data ?

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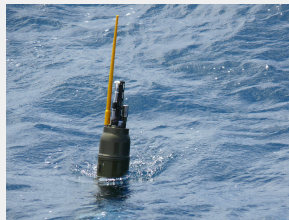
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ARGO data in the Irminger Sea

Between Sept. 2011 and Sept. 2012 :

- 574 vertical profiles from 51 different floats
- 4 floats have particularly been studied (4901163, 4901165, 4901166 and 5902298)
- 1 float with an O_2 sensor during convection

Model and satellite data

- ADT (Absolute Dynamic Topography) : AVISO (7 days ; $1/3^\circ$)
- SST, wind : ERA-Interim (12h ; $1/2^\circ$)
- Air-sea heat flux : ERA-Int. (12h ; $1/2^\circ$). NCEP (R-II) and ARPEGE (op. model) (1 day ; $1/2^\circ$)



Method of determining Mixed Layer Depths (MLD)

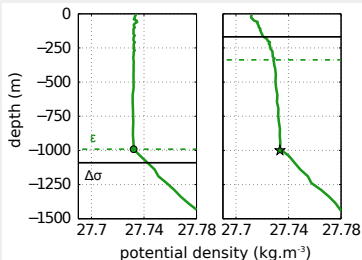
2 automatic methods

- ▶ the threshold method (de Boyer Montégut et al., 2004) : $\Delta\sigma = 0.01 \text{ kg}\cdot\text{m}^{-3}$
- ▶ the split-and-merge method (Thomson and Fine, 2003) : $\epsilon = 0.003 \text{ kg}^2\cdot\text{m}^{-6}$

Visual inspection of vertical profiles

Check results of the 2 automatic methods and identify :

- ▶ homogeneous profiles from the surface (●)
- ▶ homogeneous below a stratified layer (★)



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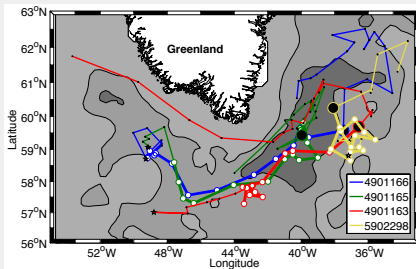
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Deepening of the mixed layers along floats trajectories

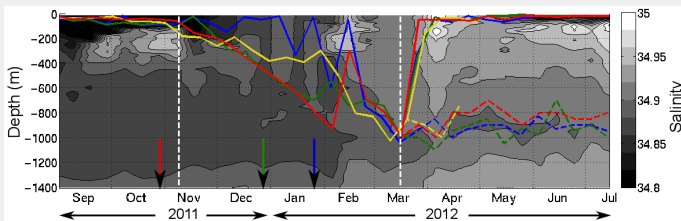


Left : floats trajectories between Sept. 2011 and Jul. 2012

- white dots : floats positions during deepening
- black dots : 1000m MLD

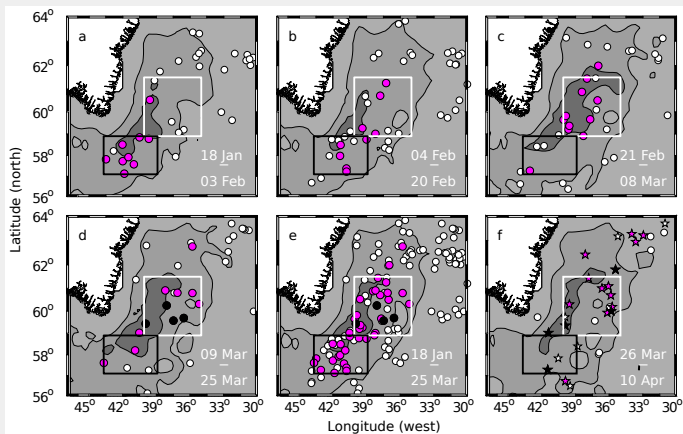
Bottom : MLD along floats trajectories

- white lines : delimit the deepening phase
- arrows : floats enter the Irminger Sea
- shade : float 4901163 salinity



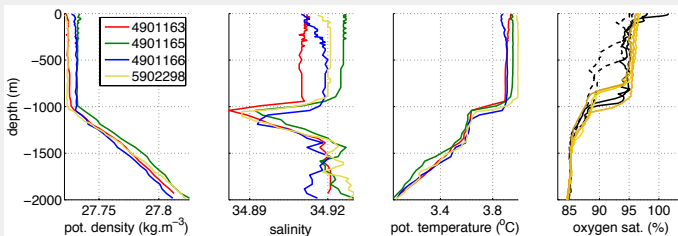


Spatial extent of the convection



Symbols : MLD (white : $< 680m$ - magenta : $\geq 680m$ - black : $\approx 1000m$)
Shade : dynamic topography (contours : -65, -55cm)
Boxes : convection areas (white : north box - black : south box)

Local formation of the convection



Right panel : oxygen profiles (every 10 days) of the 5902298 float

- black dashed : 30 Jan. ; 9 Feb.
- yellow : 19 Feb. to 20 Mar.
- black : 30 Mar. ; 9 Apr.

Preliminary conclusions :

- The convection occurs at basin scale following different stages. The deepening :
 - Y was progressive until early February
 - Y slowed in February
 - Y strongly restarted, bringing MLDs at 1000m at mid-March
- The convection formed locally in the Irminger Sea

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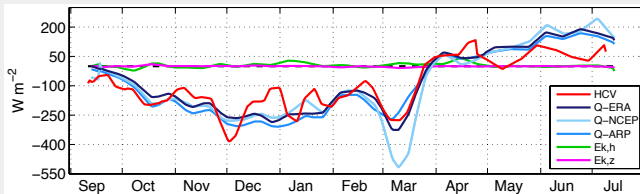
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Heat budget along floats trajectories

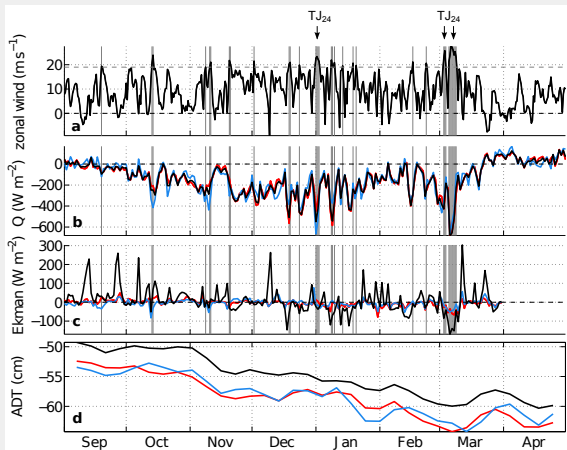
Equation of heat budget [de Boisséson et al., 2010]

$$h\delta_t \langle T \rangle = \frac{F_{net}}{\rho_0 c_p} - U_{Ek} \delta_x T - V_{Ek} \delta_y T - [\langle T \rangle - T(-h)] w_{Ek}$$



- ⤵ Heat flux explain most of the mixed layer heat variation
- ⤵ Ekman terms are weak
- ⤵ Heat losses are exceptionally strong early March

Winter 2011-2012 Greenland Tip Jets



Black curves : TJ box

Red curves : north boxes

Blue curves : south box

Tip Jet criterion

zonal wind $> 19 \text{ m.s}^{-1}$

Winter 2011-2012 inventory :

- Total : 18 TJ (Sep. - Mar.)
- 3 longer than a day
- 2 very exceptional early March



Preliminary conclusions

- Air-sea heat flux mainly explain the mixed layer heat content variation
- The heat loss events are concomitant with Greenland Tip Jets :
 - ✧ Number of Tip Jets reduced in February explain the slowing deepening during February
 - ✧ An event of strong heat loss, caused by 2 consecutive intense Tip Jets, explains the abrupt re-deepening of the mixed layers at mid-March
- Heat loss by Ekman terms are weak

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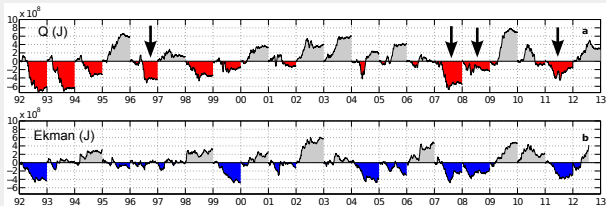
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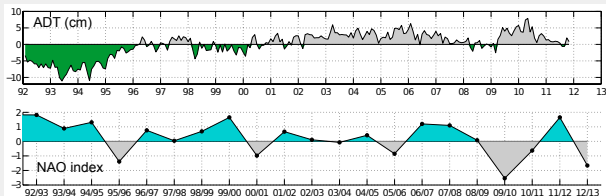
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Back to the early 1990s...



Anomalies of cumulated heat flux :
 - at the surface (top)
 - associated with Ekman transport (bottom)



- Anomalies of dyn. topography (top)
 - Hurrell PC-based NAO wintertime (DJFM) index (bottom)

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Conclusions

- Thanks to ARGO data, it is the first observation of Irminger Sea deep convection, **at basin scale** and **during the whole period of convection**
- Convection zone located east and south of Cap Farewell
- Deepening (2 months) : firstly progressive, then interrupted (Feb.), and finally strongly restarted bringing ML at 1000m. Fast restratification
- Heat budget along floats trajectories :
 - Air-sea heat flux are mainly responsible for the ML heat content variation
 - Role of Ekman terms is less important
- The TJ caused the heat loss
- Late intense TJ early March deepened the MLs up to 1000m. Without these late TJ, MLs would probably not reach 1000m and convection would have stopped in February : so, finally a local small-scale atmospheric event has influence on a larger scale oceanic event
- NAO, Ekman, Q, cyclonic circulation : indicators of deep convection for the Irminger Sea, for past years that have no observation

THANK YOU!

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