# Quality control of large Argo datasets

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## Context:

- Argo: more than 9000 profiles per month
- The **utopia**: develop an automatic tool to remove all bad data and keep all good data !!!
- The data center **approximation**: apply automatic QC to separate the problem in 2 parts
  - 1. Real time assimilation: remove all bad data (ex: ECMWF wind measurements from ships) (QC: 1)
  - 2. Delayed mode use (re-analysis): keep all possibly good data, that will need later processing (QC: 1, 2,3?)
- The **fact**: the real world is not binary and frontiers are difficult to specify in the QC continuum 1-2-3
- Major reason to the failure of automatic testing: Qualification of **individual** profiles does not exploit all available knowledge

#### Using estimation theory for QC

- Estimation theory permits combination of informations from previous knowledge (climatology and statistics) and the full dataset (neighbors): Optimal estimation (BLUE)
- Underlying assumption: zero mean and gaussian distribution of variables
- This method, can be used in delayed mode and in real time to allow for early detection of errors.
- To discriminate between the error sources, it is better applied on the basic variables (temperature and salinity on depth levels), work on elaborate variables (isotherms or isopycnals reference, dynamical height) should be introduced at a later stage, as for example
  - Wong Method
  - Comparison with altimetry

#### Analysis system: Method

#### **Objective analysis** (Optimal Interpolation)

- yo: data vector (observations)
  xa: state vector (T or S at grid points)
- Pa: covariance matrix of analyzed field
- R: data error covariance matrix
- •Co and Cao: covariance matrices of obs-obs and analyzed-obs
- •Xf: first guess for x at analyzed points
- •Yf: first guess at data point

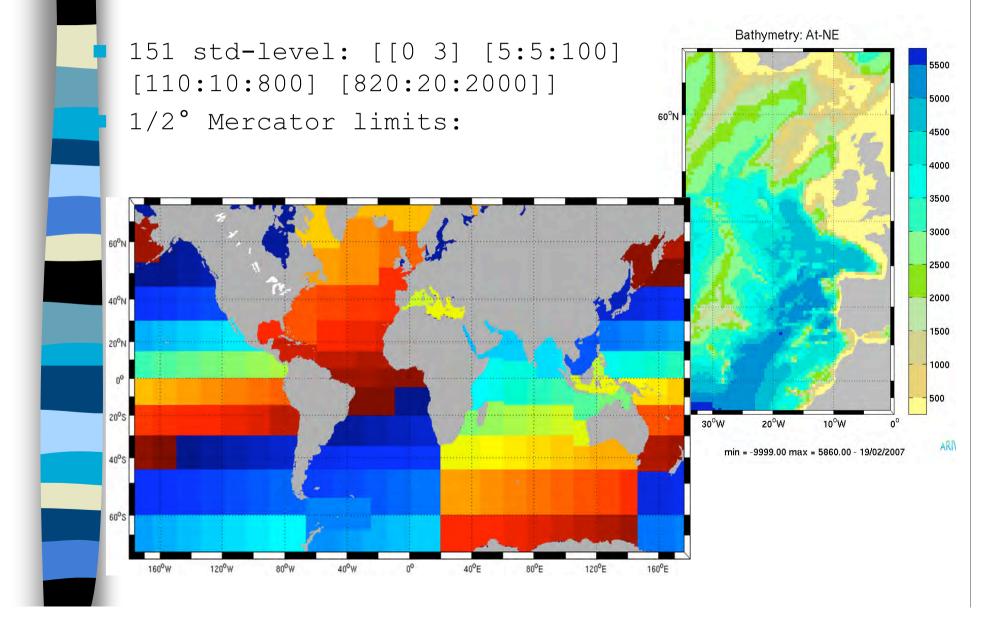
Equivalent observation/mapping matrix:

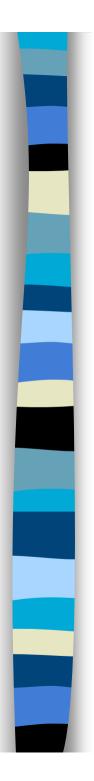
$$H^T = P^{f-1}C_{ao}$$

 $x^{a} = x^{f} + C_{ao} (C_{o} + R)^{-1} d$   $P^{a} = P^{f} - C_{ao} (C_{o} + R)^{-1} C_{ao}^{T}$   $y^{o} - y^{ao} = R (C_{o} + R)^{-1} d$ 

 $d = y^o - y^f$ 

### In Situ Analysis System : ISAS-V4 ARIVO

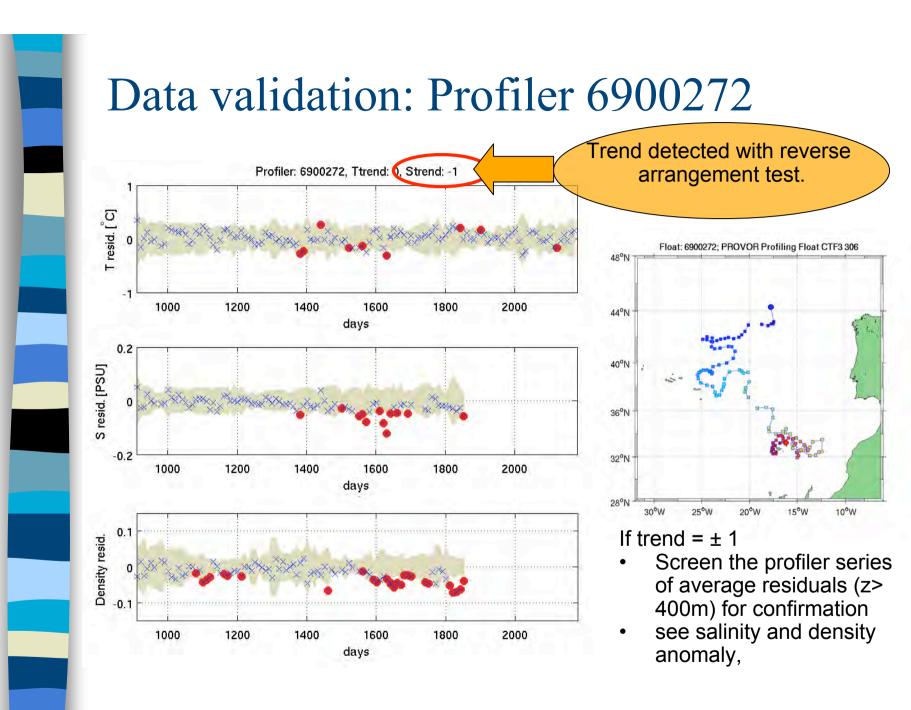




# Data validation

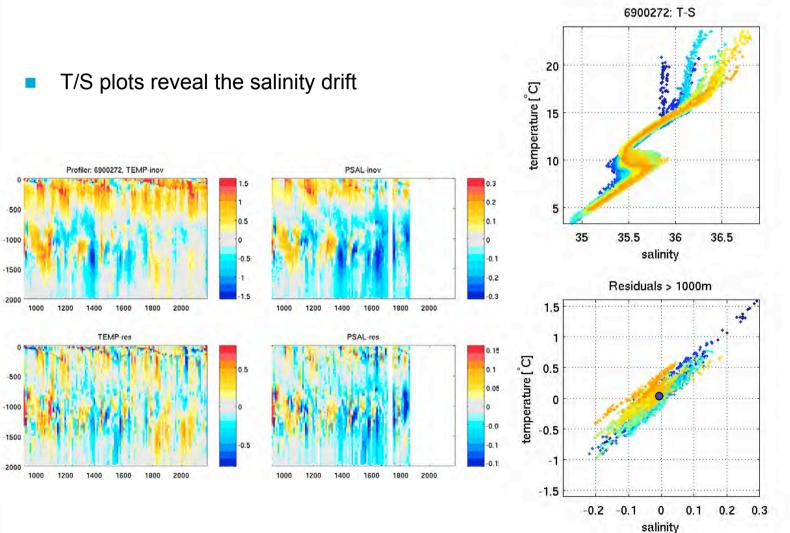
Analysis of residuals:

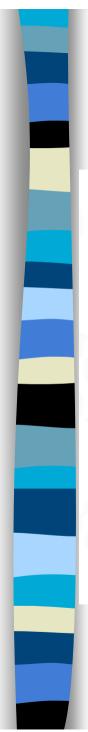
- $y_o y_a = R (C_o + R)^{-1} d$
- Applied to ARGO floats:



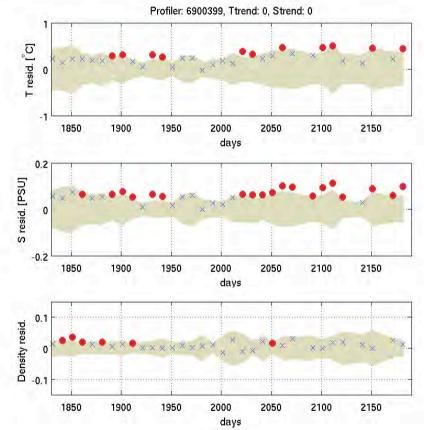


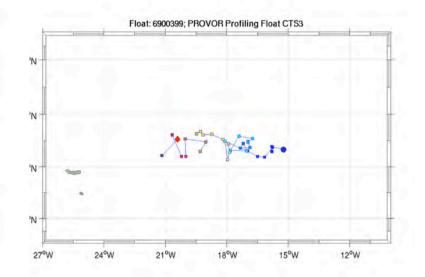
#### Data validation: salinity drift



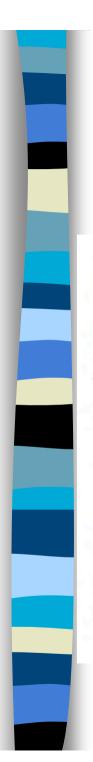


#### Data validation: PF6900299

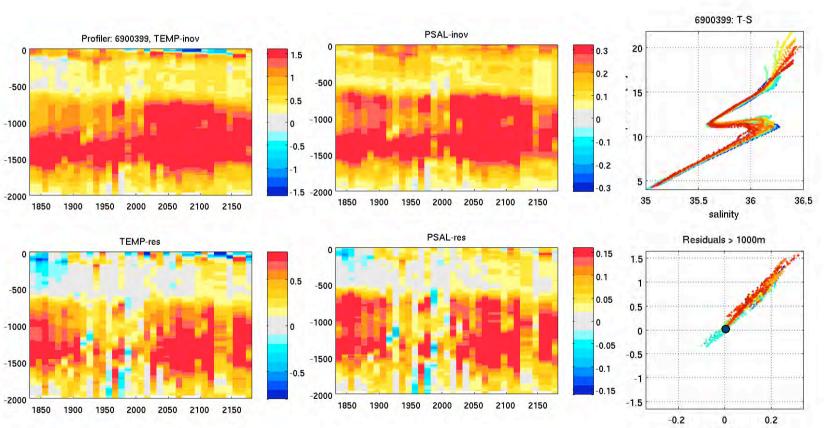


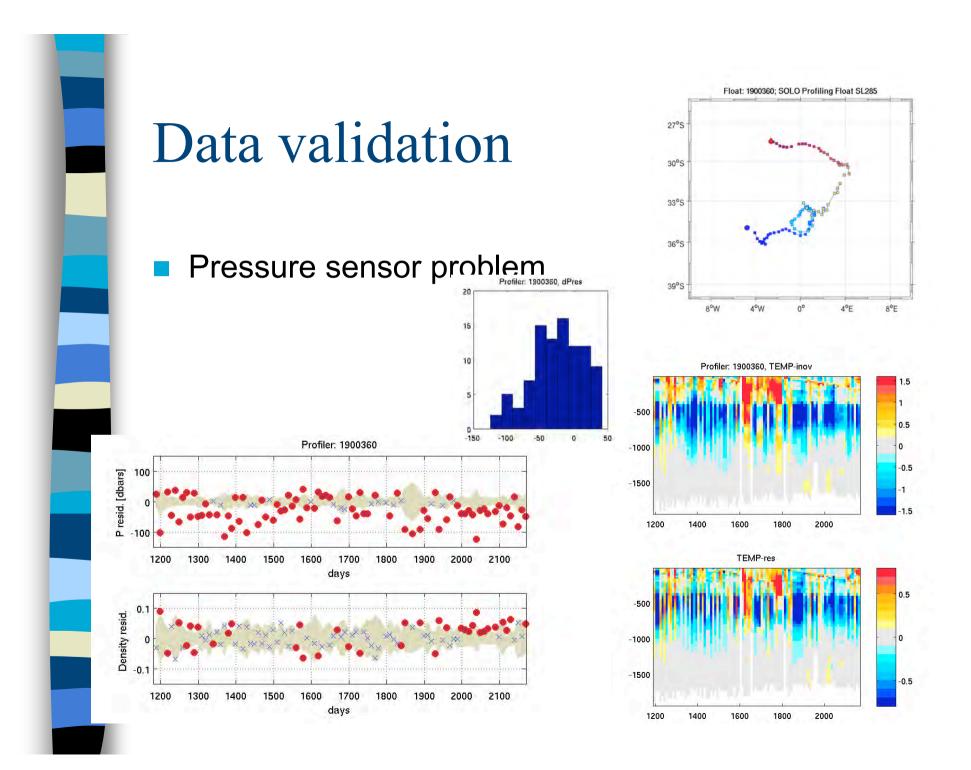


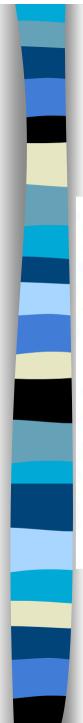
No trend detected with reverse arrangement test.



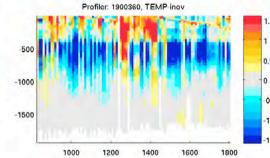
#### Profiler in a meddy

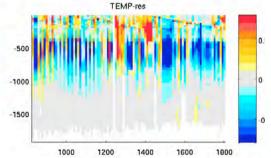


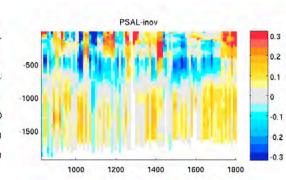


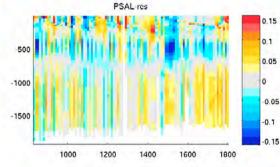


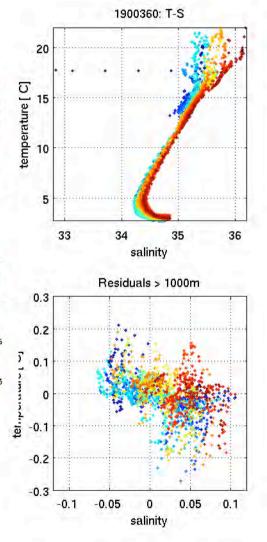
#### Data validation: Pressure problem

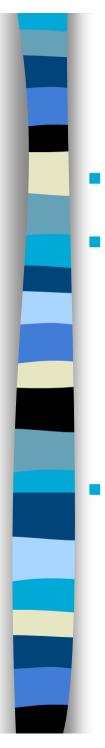






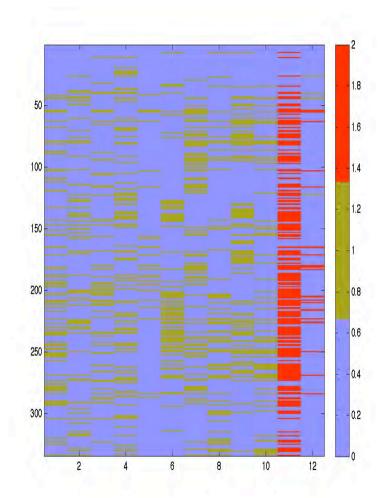






#### Converting this into numbers

- Define thresholds on various measures:
- 1 line = 1 profiler, 1 column = 1 test
  - Col 1, 2: T-offset, Tdrift
  - Col 3, 4: S Offset, Sdrift
  - Col: 5, 6: Density offset and STD
  - Col 7, 8, 9: Pressure offset, STD and Pout
  - Define combined indices:
    - Col 11: Pressure pb.
    - Col 12: opposite T-S drift or offset (expression of T sensor error)



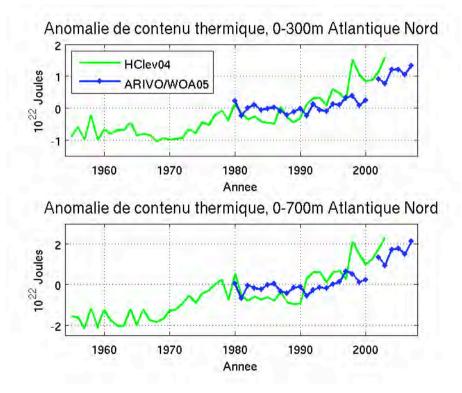


# Conclusion

- Obtaining a perfect dataset will need many iterations and requires to alternate simple automatic QC, global statistical coherence tests and single profiler screening
- A first pass on the 2002-2007 dataset already gives robust results

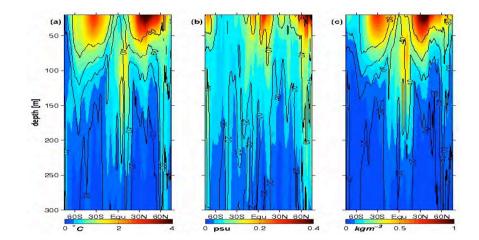


## Heat content anomaly



- Blue: anomaly of ARIVO analysis relative to WOA05
- Green:Levitus 2004 anomaly (fitted on ARIVO series over 1980-2000)

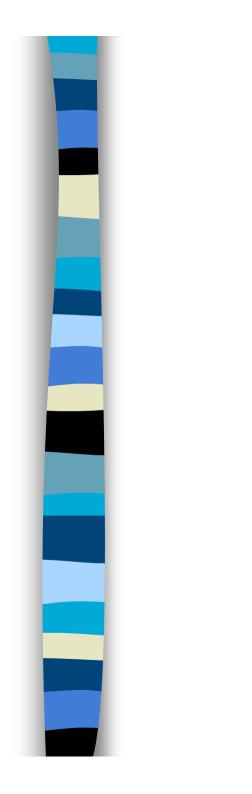
#### Global ocean variability (2003-2007) (K. Von Schuckmann) Seasonal cycle: The 12 month harmonic (extracted from the 5 years series) show penetration of the seasonal cycle deeper than 200m.



#### Océan gobal Pacifique 300 600 900 1200 1500 1800 Atlantique Indien 60N Equ 30N 60N 305 Eau 30N 300 600 900 1200 1500 1800 60N 60S 60S 305 60N 305 Equ 30N Equ 30N 0.05 0.1 psu 0

#### Intra-saisional to interannual variability:

Zonal mean of the standard deviation show the strong deep variability of the North Atlantic (mid and high latitudes) and of the southern ocean (south of 30S)





## Annex



#### Covariances scales

$$C_{o} = C_{o}(dx_{o}, dy_{o}, dt_{o}) = \sum \sigma_{i}^{2} e^{-\left[\frac{dx_{o}^{2}}{2L_{x}c^{2}} + \frac{dy_{o}^{2}}{2L_{y}c^{2}} + \frac{dy_{o}^{2}}{2L_{y}c^{2}}\right]},$$

i = 1; Lx=Ly=300km; Lt = 21 days i = 2; Lx=Ly=2\*resol<4\*Rossby radius)<300km; Lt = 21 days

